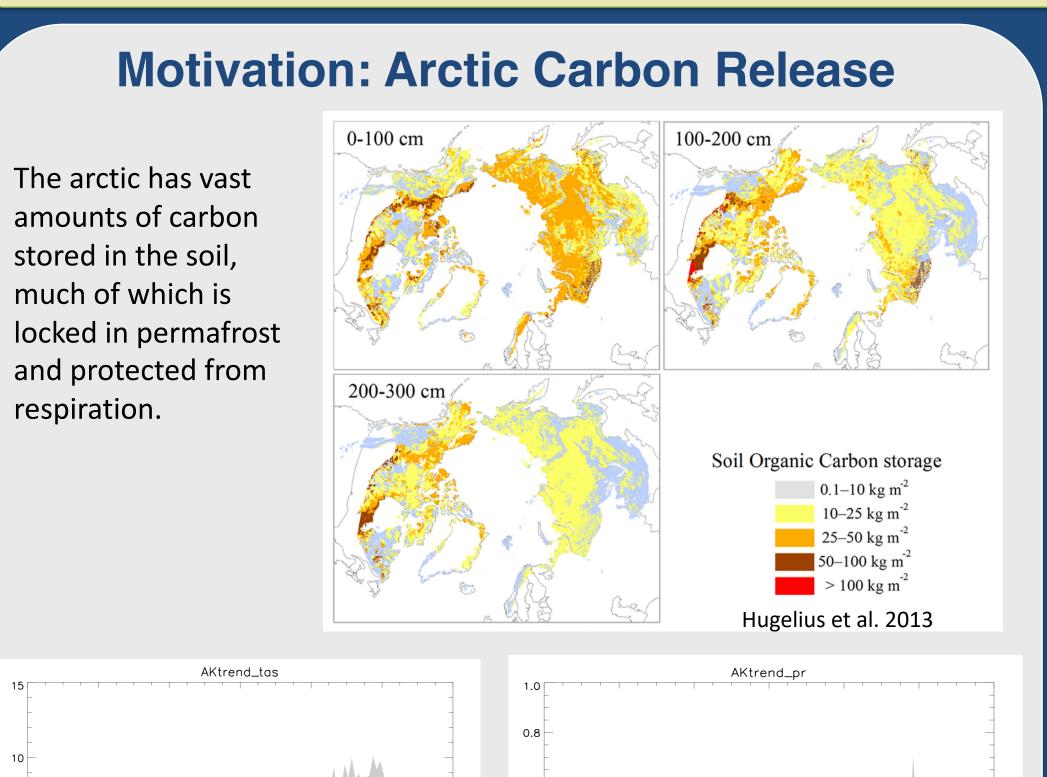
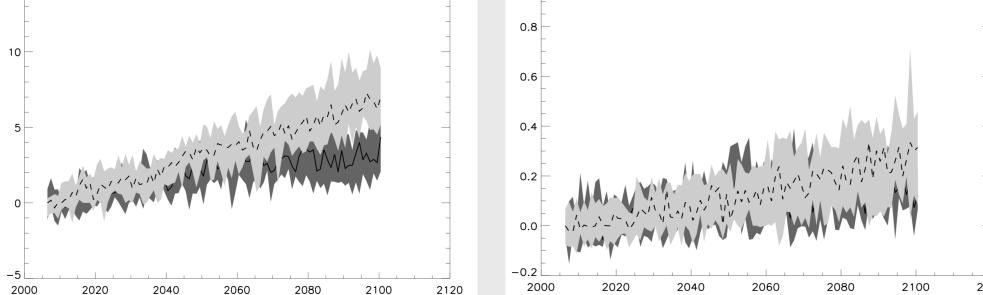
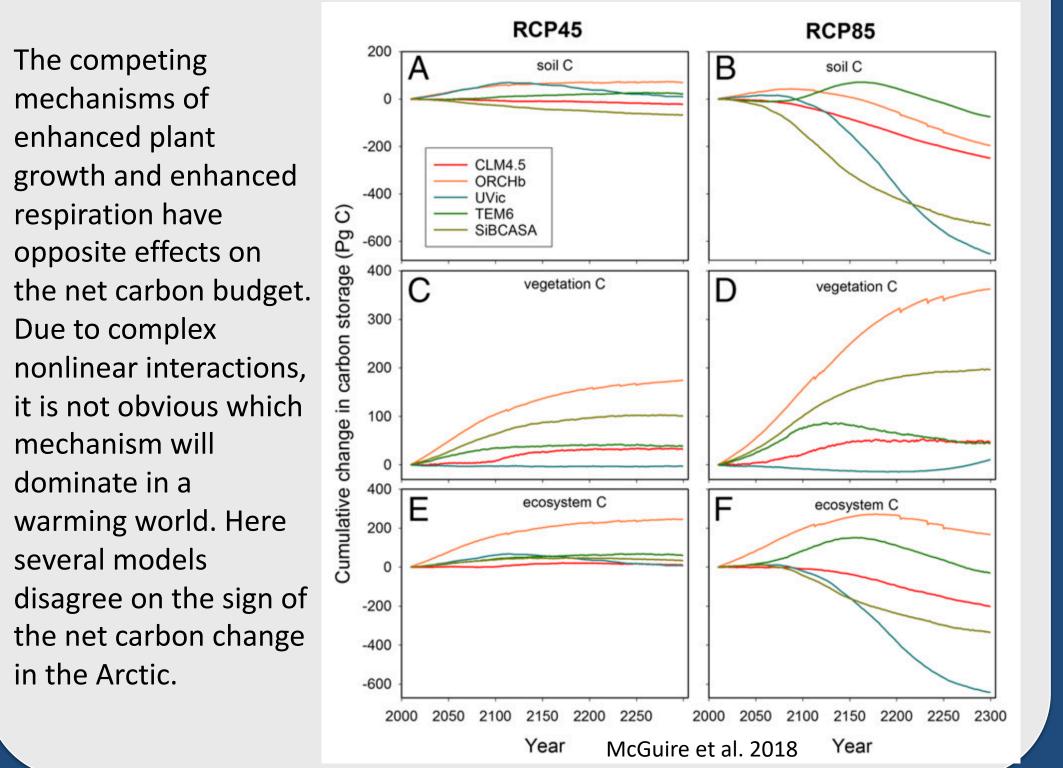


Erik Larson<sup>1</sup>, Bill Munger<sup>1</sup>, Roisin Commane<sup>1,2</sup>, Luke Schiferl<sup>1,2</sup>, Steve Wofsy<sup>1</sup>, Eugenie Euskirchen<sup>3</sup>, Donatella Zona<sup>4</sup>, Takeshi Ise<sup>5</sup>, Paul Moorcroft<sup>1</sup> <sup>1</sup>Harvard University, <sup>2</sup>Columbia University, <sup>3</sup>University of Alaska, <sup>4</sup>Sandiego State University, <sup>5</sup>Kyoto University contact: erik\_larson@fas.harvard.edu

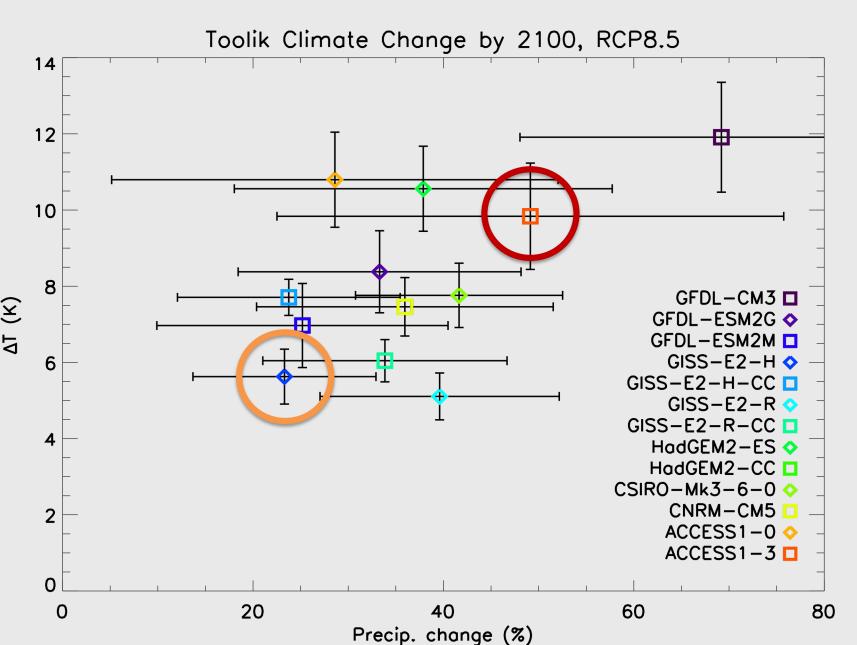




However, this carbon may become accessible as temperature (K) and precipitation (fraction) change in the future. Here are the RCP8.5 and RCP4.5 scenarios for the Alaskan tundra.



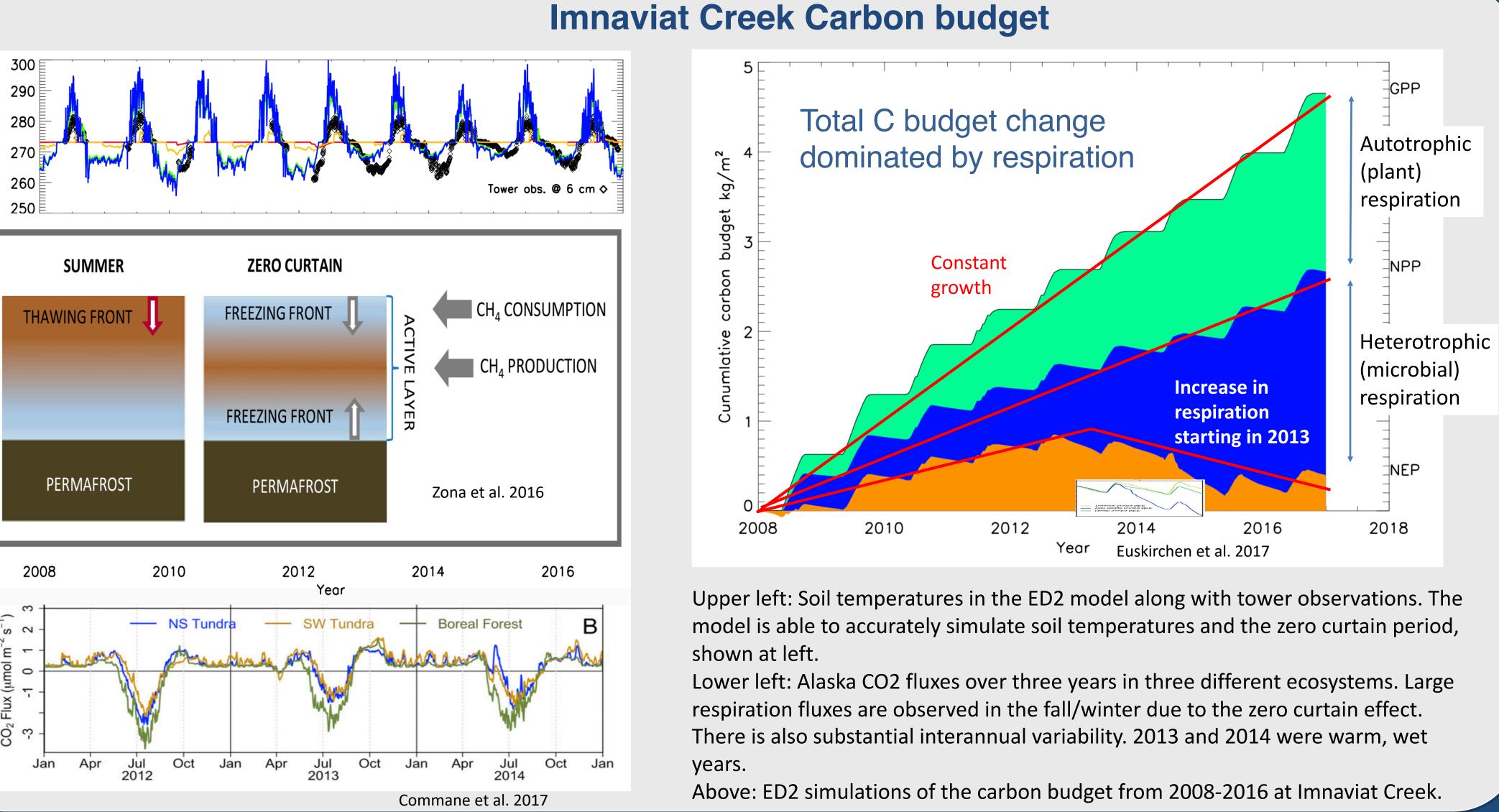
# **RCP8.5** Warm and Hot scenarios



We simulate two year 2100 climates based on CMIP5 RCP8.5 models that span the range of climate change. We apply the changes in the monthly climate variables to 3 hourly tower meteorological data to create the input meteorology. The model is driven with specific humidity, precipitation, short and long wave surface radiation, and temperature.

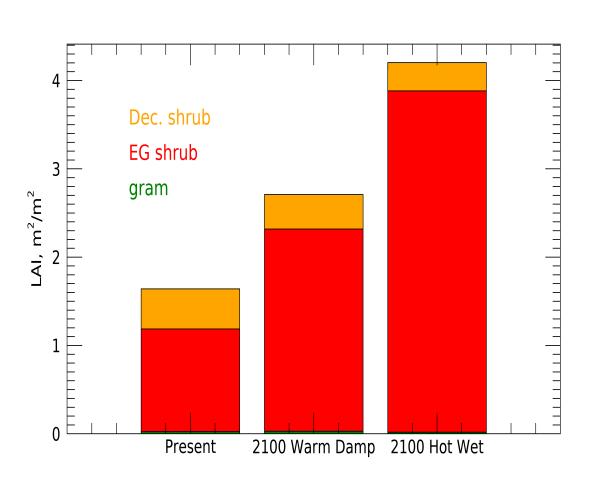
The contour in the relative soil moisture panels is the soil liquid fraction, and therefore tracks the temperature. Increased plant transpiration substantially dries the soil column in the growing season of the hot scenario, despite a 50% increase in precipitation.

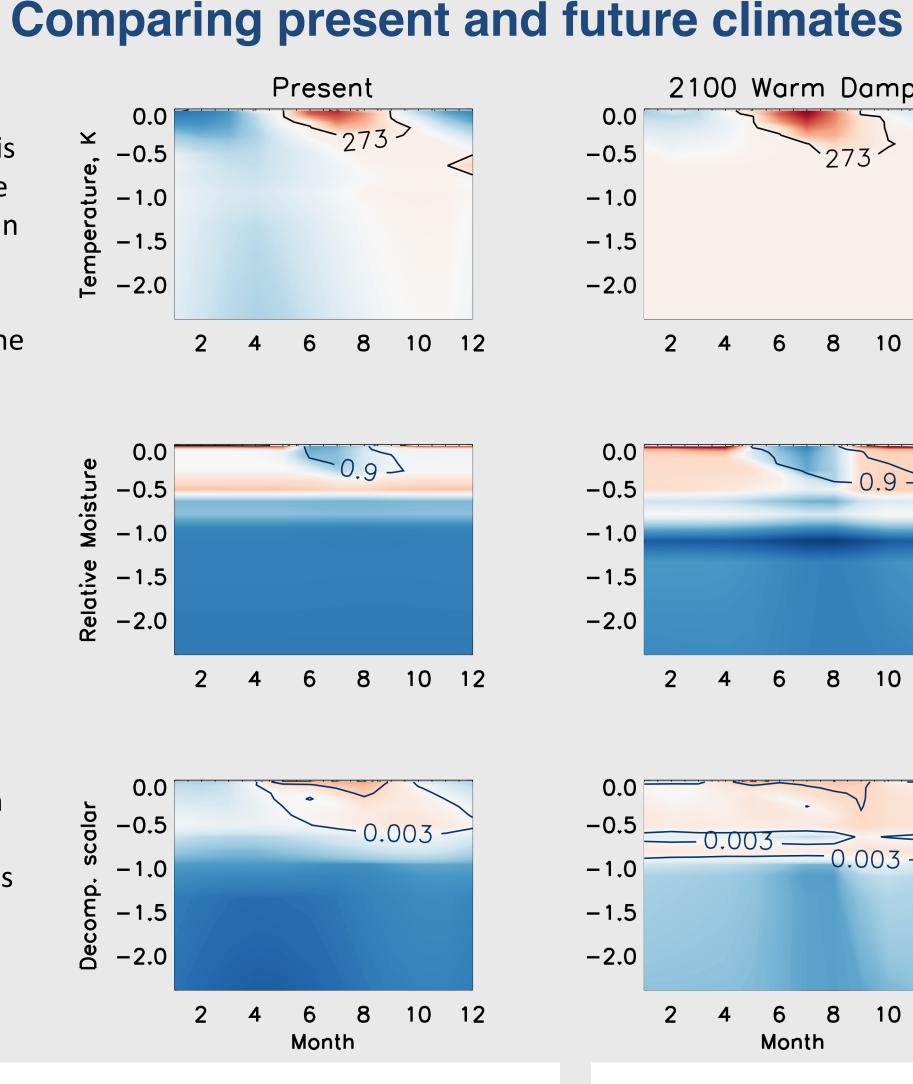
# Partitioning changes to the Arctic carbon budget in future RCP scenarios

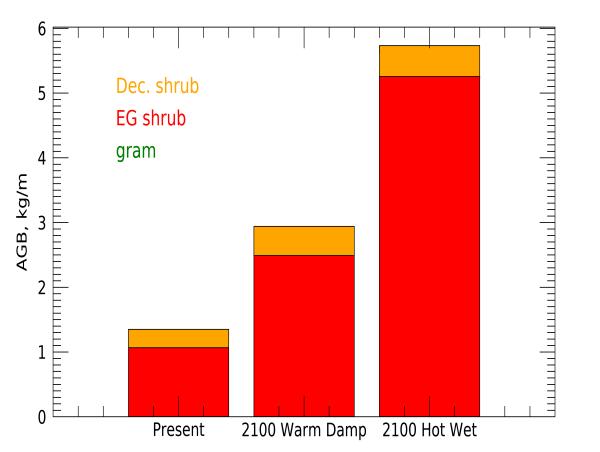


Present and future soil temperature, moisture, and respiration scalar as a function of depth and time. This is smoothed over 10 years for 3 climate scenarios, the present, 2100 "warm damp", and 2100 "warm wet". In warming climates, the whole soil column warms, and the active layer deepens. At some point between our two warming scenarios we cross a threshold where the entire soil column melts (at least down to 2.4 m).

Interestingly, in the warmest scenario, the respiration increases dramatically, as expected, but not in the summer. Most of the increase is in the fall/winter. This is due to increased plant transpiration drying the soil column in the growing season and decreasing the heterotrophic respiration.



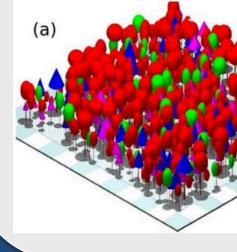


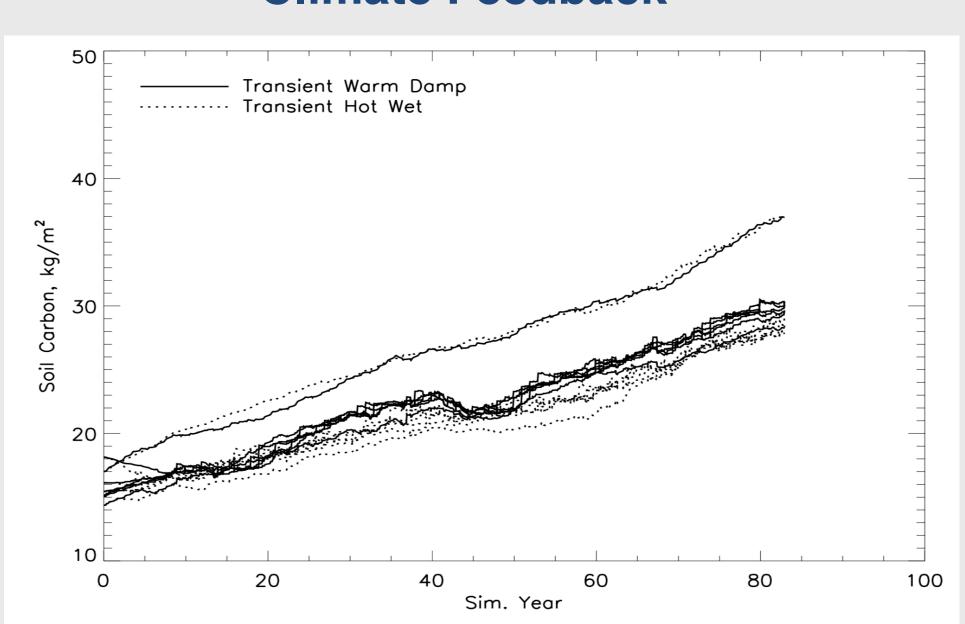


Above Left: Equilibrium LAI and AGB in present and future climates by PFT.

Above Right: Carbon fluxes in present and future climates. Year 2100 carbon fluxes are much larger and carbon is being stored in the soil.

- Atmospheric grid ce carbon Patch 4 Patch (harvested Fluxes: W. H. ( patches interact, Disturbance:  $\lambda_{F}$ the use of light, Dispersal and recruitment: moisture and nutrients. Stem growth: gs N(y,t) Medvigy et al. 2009 Gap y
- biosphere model internal energy, and Sites are driven by met data and characterized by unique geophysical morphology including soil type and hydrology representations of ecosystems. Plants in compete and facilitate
- Mechanistic terrestrial • Conserves water, Patches are statistically

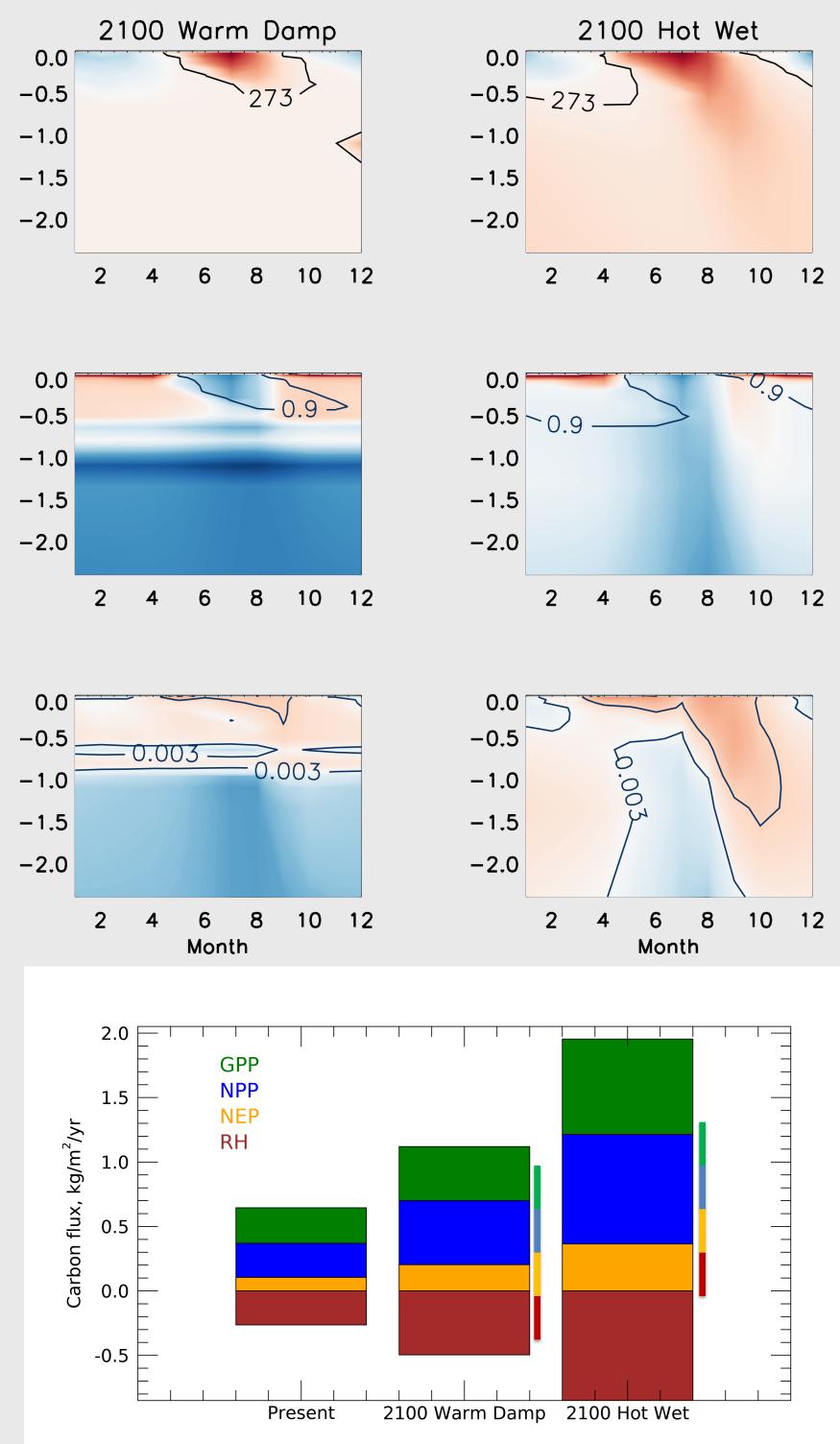




### **Acknowledgements and References**

Medvigy et al. 2009. Mechanistic scaling of ecosystem function and dynamics in space and time: Ecosystem Demography model version 2. JGR Biogeo, https://doi.org/10.1029/2008JG000812 Commane et al. 2017. Carbon dioxide sources from Alaska driven by increasing early winter respiration from Arctic tundra. PNAS, 114 (21), 5361-5366 Euskirchen et al. 2017. Long-term release of carbon dioxide from arctic tundra ecosystems in Alaska. Ecosystems, 1-15 McGuire et al. 2018. Dependence of the evolution of carbon dynamics in the northern permafrost region on the trajectory of climate change. PNAS, 115, (15), 3882–3887 Zona et al. 2016. Cold season emissions dominate the Arctic tundra methane budget. PNAS, 113 (1), 40-45

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## **ED2 Description**

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# **Climate Feedback**

On average, 15 kg C/m<sup>2</sup> is expected to be added to the Imnaviat Creek soil by 2100. This corresponds to a net increase in soil carbon in the Alaskan Tundra of about 2.3 Pg, with another 0.5 Pg in plant biomass. We are not seeing, by 2100 at least, a net loss of carbon from the ecosystem.

2.8 Pg of C uptake by the Alaskan Tundra ecosystem by 2100 results in a meager change in atmospheric  $CO_2$  of -1.3 ppmv.

Over the next 100 years, the Alaskan tundra will take up roughly the equivalent of the United States anthropogenic CO<sub>2</sub> emissions in this year.